

鄂尔多斯盆地东部山西组页岩气成藏特征及勘探对策

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内容提要:我国海陆过渡相页岩气资源量约占全国资源总量的三分之一,勘探开发前景广阔。鄂尔多斯盆地过渡相页岩分布范围广、厚度大,受到广泛关注,勘探开发一直未能获得突破。为查明过渡相页岩地质特征与海相页岩储层特征及成藏规律的差异,本文以鄂尔多斯盆地东部山西组页岩为例,通过总有机碳、镜质组反射率、矿物组成、微观孔隙和含气量分析开展页岩气生成、储集和成藏过程的对比研究,分析认为山西组页岩成藏过程有其特殊性,其干酪根来源于Ⅲ型高等植物,液态烃产量很少,缺乏气泡成孔过程,有机质纳米孔数量少,砂泥煤互层的沉积特征排烃效率高,影响页岩含气性。山西组页岩气形成后,大部分页岩气已运移到附近砂岩中形成致密气,致密砂岩、煤层和页岩呈三明治式成藏。预测研究区内东南部为有利区域。在勘探山西组页岩气时考虑煤系“含气带”,通过立体勘探寻找有利的富气层段。

关键词:鄂尔多斯;山西组;页岩;有机质纳米孔

我国页岩气勘探开发工作在最近几年取得了较快发展,海相页岩气勘探开发关键技术已经取得重大突破,在四川盆地长宁、威远、涪陵和云南昭通地区建成了国家级页岩气示范区。除海相地层外,海陆过渡相地层是重要领域。据统计,我国海陆过渡相页岩气可采资源 $8.97 \times 10^{12} \text{ m}^3$,鄂尔多斯盆地占一半以上(Ross and Bustin, 2007; Yu Bingsong, 2012; Gong Meilin et al., 2013; Dong Dazhong et al., 2016)。与海相地层相比,海陆过渡相利于形成页岩气、煤层气和砂岩气等多种类型天然气叠置气藏(Wang Shejiao et al., 2011; Feng Ziqi et al., 2013; Yan Deyu et al., 2015; Sun Zepeng et al., 2017)。过渡相地层往往呈现页岩频繁与煤、致密砂岩互层,纵横向变化快,单层厚度小,累计厚度大的特征(例如, Yan Deyu et al., 2013, 2015; Fu Jinhua et al., 2013; Zhao Keying and Guo Shaobin, 2015; Lan Chaoli et al., 2016; Zhou Shuai et al., 2016)。海陆过渡相页岩有机质含量变化较大,有机质类型以混合型-腐殖型为主(例如, Zhang Jizhen et al.,

2015; Lan Chaoli et al., 2016; Tang Xuan et al., 2016; Xiong Fengyang et al., 2017; Yang Chao et al., 2017);孔隙类型主要有粒间孔、晶间孔、溶蚀孔、微裂缝和有机质孔隙,其中有机质孔发育程度差异较大(例如, Zhao Keying and Guo Shaobin, 2015; Zhang Jizhen et al., 2015; Xiong Fengyang et al., 2017; Zhang Qin et al., 2017);黏土矿物含量较高,可压裂性较海相地层页岩差(例如, Fu Jinhua et al., 2013; Lan Chaoli et al., 2016)。目前,国内针对海陆过渡相页岩气藏的研究主要集中在四川盆地周缘龙潭组(Liu Guangxiang et al., 2015; Zhang Jizhen et al., 2015; Meng Fanyang et al., 2017)和鄂尔多斯盆地山西组、太原组和本溪组(例如, Fu Jinhua et al., 2013; Guo Wei et al., 2015; Zhao Keying and Guo Shaobin, 2015; Lan Chaoli et al., 2016; Tang Xuan et al., 2016; Xiong Fengyang et al., 2017; Zhang Qin et al., 2017)。一些生产公司开展了海陆过渡相页岩气勘探开发试验,在一些探井获得页岩气流,但是还没有获得商业突破。本文

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Liu Honglin, Wang Huaichang, Zhang Hui, Zhao Weibo, Liu Yan, Liu Dexun. 2020. Geological characteristics and exploration countermeasures of shale gas in the Shanxi Formation of the Ordos basin. Acta Geologica Sinica, 94(3): 905~915.

以鄂尔多斯盆地山西组页岩为研究对象,通过对露头 and 岩芯样品进行总有机碳(TOC)、镜质组反射率(R_o)、全岩矿物组成、微观孔隙类型和含气量分析,分析山西组页岩储层地质特征,在此基础上探讨山西组页岩气藏成藏特殊性,并提出经济勘探开发山西组页岩的勘探策略,对于加快海陆过渡相页岩气勘探开发具有重要意义。

1 区域地质概况

鄂尔多斯盆地属于华北克拉通的一部分,现今基本构造格局形成于燕山期,发展完善于喜马拉雅期,整个盆地构造面貌总体为一东翼缓而长、西翼陡而短、南北翘起的不对称大向斜。根据盆地基底特征、构造形态,结合沉积及油气分布等特点,分为 6 个一级构造单元:渭北挠褶带、西缘冲断推覆构造带、天环凹陷、陕北斜坡、晋西挠褶带、伊盟隆起(Zhao Keying and Guo Shaobin, 2015; Sun Zepeng et al., 2017)。鄂尔多斯盆地晚古生代经历了陆表海、近海湖盆以及内陆拗陷湖盆等古地理演化过程(Yan Deyu et al., 2013; Zhou Shuai et al., 2016)。研究区位于鄂尔多斯盆地东部(图 1)。

研究区过渡相地层主要发育在中石炭统本溪组、上石炭统太原组和下二叠统山西组。其中,本溪组以障壁岛-潮坪-泻湖为主;太原组发育辫状河-三角洲-泻湖-潮坪-障壁岛-浅海陆棚碳酸盐岩沉积;山西组以三角洲-潮坪沉积为主。总体来看,本溪组和太原组页岩欠发育,而山西组页岩普遍厚度大,为本文研究目的层段。

山西组根据岩性与电性特征划分为山一段和山二段。山一段以灰色砂岩与深灰色页岩互层为主,局部发育薄层煤线;山二段以黑色页岩与灰色砂岩为主,夹 3~5 层煤层,煤厚度普遍 2~7m 不等,其内部可划分为山₂¹、山₂² 和山₂³ 三个砂组(图 2)。山西组页岩普遍具有“四高两低”特征,即高伽马(GR)、高电阻率(RT)、高声波时差(AC)、高补偿中子(CNL)、低密度(DEN)和低光电效应(图 2)。

2 样品和分析方法

本次研究在岩芯描述和露头考察基础上,共采集代表性露头 and 岩芯样品 45 个。页岩 TOC 使用 LECO-CS230 碳硫分析仪测试获得,依据标准为国标 GB/T 19145—2003;R_o 使用 Axio Scope A1 显微光度计测定,通过测定沥青反射率计算得到等效镜质组反射率,测试依据分别为中国石油天然气行

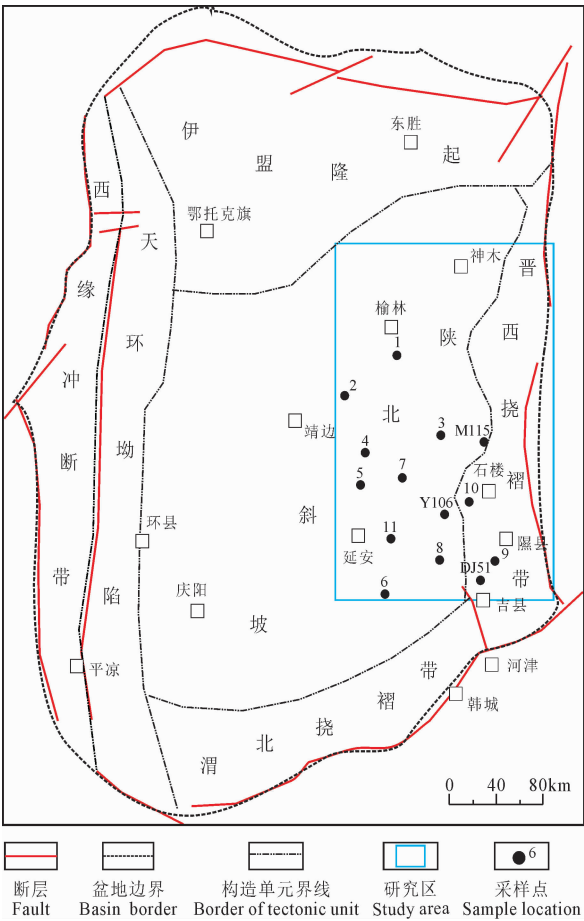


图 1 鄂尔多斯盆地构造分区图
Fig. 1 Tectonic subdivision of Ordos basin

业标准 SY/T5124—2012 和 SY/T5125—2014;矿物组成使用 D/Max-2500PC X 射线衍射仪进行测定,测试依据为 SY/T5163-2010;微观孔隙采用氩离子抛光仪 Gatan Model 697 进行抛光,然后使用 FEI Heilos 650F 型扫描电镜进行微观孔隙观察;孔隙度用氦孔仪进行测定,渗透率采用脉冲方法获得。所有试验均在中国石油勘探开发研究院完成。

由于实测 TOC 含量样品较少,为更全面评价研究区山西组 TOC 含量分布,采取利用井实测 TOC 含量标定测井资料,再根据测井资料预测 TOC 含量的方法来评价研究区 TOC 含量的平面分布。由于富含有机质泥、页岩在测井上具有高 RT、高 AC、高 GR、低 DEN 特征,综合利用这 4 条曲线可定量预测 TOC 含量(Xu Xiaohong et al., 1998; Wang Guiwen et al., 2002; Kinley et al., 2008; Yang Taotao et al., 2013)。根据井实测的 TOC 值和实测样品深度所对应的 AC、DEN、GR 及 RT 测井拟合的 TOC 含量计算方程为:

$$TOC=0.02268\times AC+4.776595\times \log RT+$$

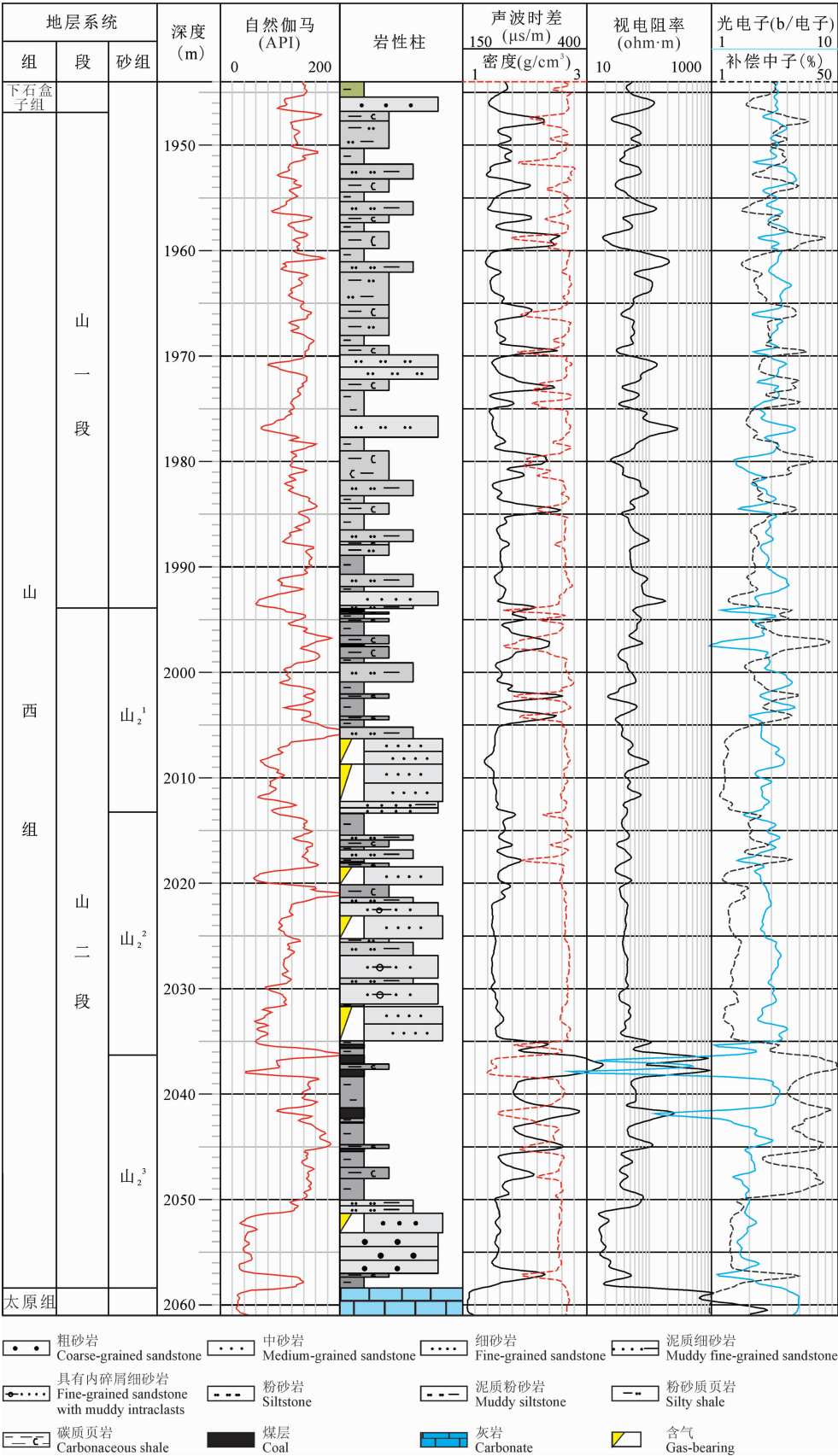


图 2 鄂尔多斯盆地东部山西组地层柱状图(Y106 井)

Fig. 2 Stratigraphic column of Shanxi Formation in eastern Ordos basin (well Y106)

$1.078298 \times \text{DEN} - 0.03825 \times \text{GR} - 11.8104$ (1)

式中:AC 为实测的声波时差($\mu\text{s}/\text{m}$);RT 为实测的电阻率($\Omega \cdot \text{m}$);DEN 为实测的密度(kg/m^3);GR 为实测的自然伽马(API)。

含气量根据现场解析实测获得。考虑到实测含气量数据较少,对缺少实测含气量数据的井点,根据公式(1)计算的 TOC,利用 Guo Wei et al. (2015)建立的 TOC 与含气量相关性关系,计算各井目的层含气量,用以预测含气量的平面变化。

3 山西组页岩储层地质特征

3.1 有机碳特征

研究区山西组山一段 TOC 含量 0.04%~19.4%,山二段页岩 TOC 含量 0.07%~16.5%,山西组页岩 TOC 峰值 1.0%~3.0%(图 3)。相比之下,山二段 TOC 更高。平面上,山西组 TOC 含量从东向西,总体有减小的趋势,其中东部存在两个高 TOC 含量区域,TOC 含量超过 3%,即府谷—神木—佳县和绥德—延川以东区块(图 4)。

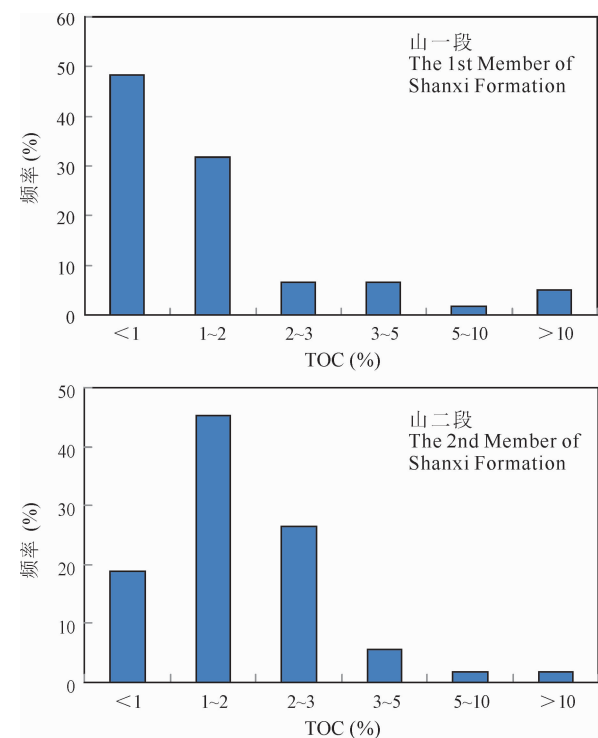


图 3 鄂尔多斯盆地东部山西组页岩 TOC 分布直方图

Fig. 3 Histogram map showing TOC of Shanxi shales in eastern Ordos basin

3.2 页岩平面分布特征

鄂尔多斯盆地东部在山西期形成了多套(碳质)页岩-灰色砂岩-(碳质)页岩-煤的三角洲-河流相沉

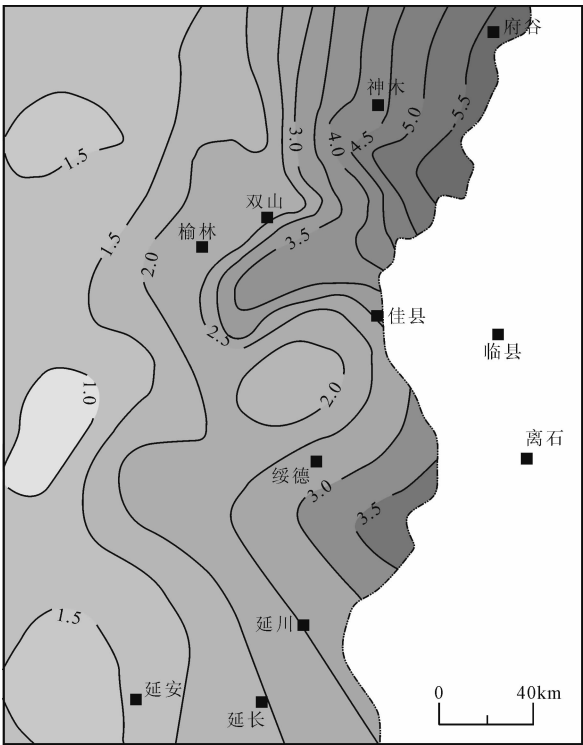


图 4 鄂尔多斯盆地东部山西组 TOC(%)等值线图

Fig. 4 TOC (%) isopach of Shanxi shales in eastern Ordos basin

积旋回。其中页岩普遍比较发育,一般 3~7 层。山一段页岩厚度 9.7~51.5m,平均 24.6m;山二段页岩厚度 21.4~92.3m,平均 41.2m。平面上,山西组页岩厚度中心分布在中部的榆林南—绥德西、东部的绥德—佳县以及南部的延安北—延川北,其中绥德—佳县地区最大厚度超过 90m(图 5)。

3.3 页岩储层微观特征

扫描电镜揭示研究区山西组页岩储层的孔隙类型主要有粒间孔、粒内孔、有机质孔和晶间孔(图 6)。从孔隙分布的比例来看,存在 20%~30%左右的有机质孔,孔隙直径一般小于 100nm,主要分布在 2~50nm 之间。孔隙形态呈开放状态,以两端开口的圆筒孔、四边开放的平行板孔(圆锥、圆柱、平板)及细颈瓶状(墨水瓶)孔等开放性孔为主(Fan Wentian, 2019; Cao Taotao et al., 2016; Zhang Yan et al., 2017; Misch et al., 2016)。页岩孔隙度主要在 0.25%~4.85%之间,平均 2.3%,实测渗透率主要在 $0.01 \times 10^{-3} \sim 0.1 \times 10^{-3} \mu\text{m}^2$ 之间,平均 $0.04 \times 10^{-3} \mu\text{m}^2$,且孔隙度与渗透率具有正相关关系。

3.4 页岩矿物组成

山西组山二段黑色页岩石英含量 8.2%~40.6%,黏土矿物含量 55.3%~70%,其余次要矿

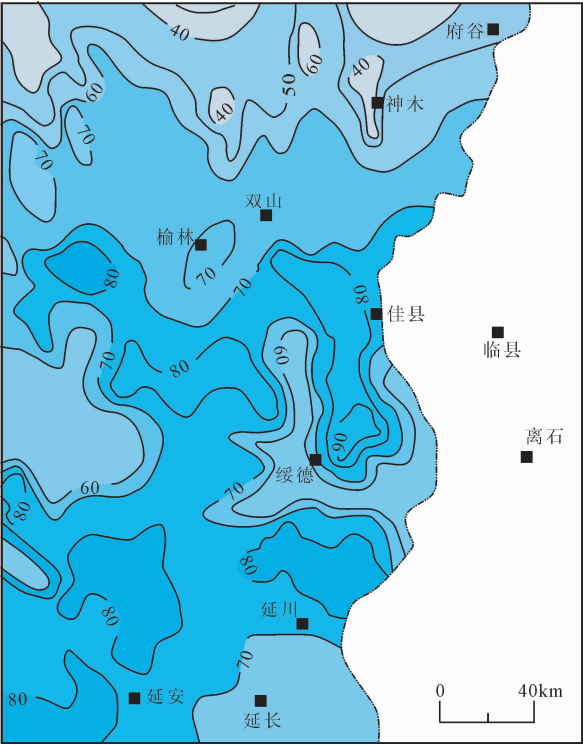


图 5 鄂尔多斯盆地东部山西组页岩等厚图(m)
Fig. 5 Thickness isopach (m) of Shanxi shales in eastern Ordos basin

物见黄铁矿、白云石、长石等。脆性矿物含量 9.5%~62.6%，平均 41.3%(图 7)。整体上，与美国和中国四川盆地海相页岩相比(Nie Haikuan et al. , 2009; Zou Caineng et al. , 2010; Li Jianzhong et al. , 2012)，中国鄂尔多斯盆地东部山西组页岩脆性矿物较低而黏土矿物含量较高。

3.5 页岩成熟度

根据实测数据编制的鄂尔多斯盆地东部山西组

页岩 R_o 。等值线图显示(图 8)，目的层页岩 R_o 介于 1.2%~2.6% 之间，平面上自北东向南西升高，与页岩埋深变化趋势具有良好的相关性。其中，沿绥德—榆林一线的东北部大部分地区小于 2.0%，该线南侧较高，一般超过 2.0%，部分地区达到 2.6%，反映研究区页岩普遍达到了生气高峰(Schenk et al. , 1997; Fu Jinhua et al. , 2013)。

3.6 页岩含气性

研究区山西组页岩含气量总体较低，普遍低于 2.0 m³/t，多介于 0.5~2.0 m³/t 之间。平面上，页岩含气量总体自北东向南西升高，高于 2.0 m³/t 的区域分布在绥德—延川—延长以西至延安以东的区域(图 9)。这种含气量的平面变化趋势显示了与 TOC 和 R_o 平面分布(图 4 和图 5)具有较好的相关性。与海相页岩含气量(Nie Haikuan et al. , 2009; Wang Shiqian et al. , 2009; Zou Caineng et al. , 2011; Li Jianzhong et al. , 2012)相比，山西组页岩含气量普遍较低。

4 山西组页岩气藏成藏规律

山西组页岩及其共生砂岩与煤形成在海陆过渡相环境。这导致山西组页岩成藏条件与海相页岩成藏条件有两个显著不同的特征。

4.1 有机质类型及其成熟过程

海相页岩中的有机质来源主要为浮游生物，包括各种藻类、浮游动物，富含蛋白质、类脂化合物及部分碳水化合物，含有相对中等分子量的正构烷烃(C₁₂~C₂₀)、正构脂肪酸、丰富的甾类(C₂₇ 胆甾烷)、类异戊二烯烷烃(C₁₄~C₂₀)，甾萜比通常大于 1 (Yang Cai et al. , 2017; Zhang Qin et al. , 2018)。

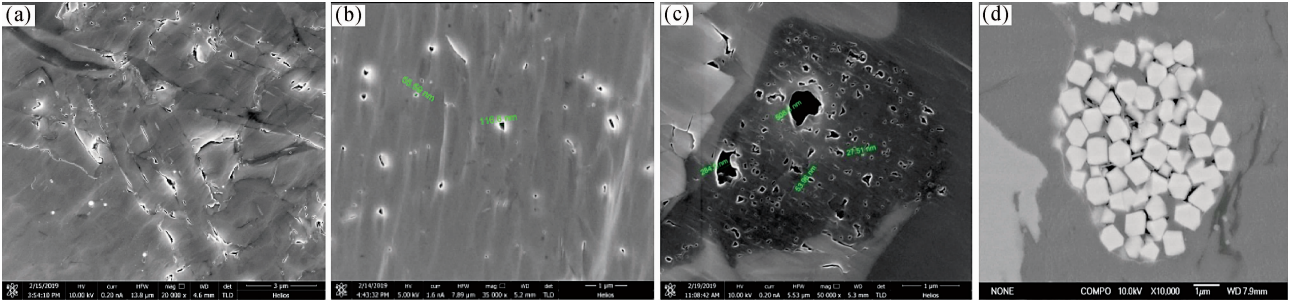


图 6 鄂尔多斯盆地东部山西组页岩孔隙类型扫描电镜图

Fig. 6 Typical FIB-SEM micrographs showing pore types of Shanxi shales in eastern Ordos basin

(a)—Well M115, intergranular pores, the Second Member of the Shanxi Formation; (b)—Well Y106, pores within clastic grain, the Second Member of the Shanxi Formation; (c)—Well DJ51, pores within organic matter, the Second Member of the Shanxi Formation; (d)—Well DJ51, intercrystal pores among pyrites, the Second Member of the Shanxi Formation

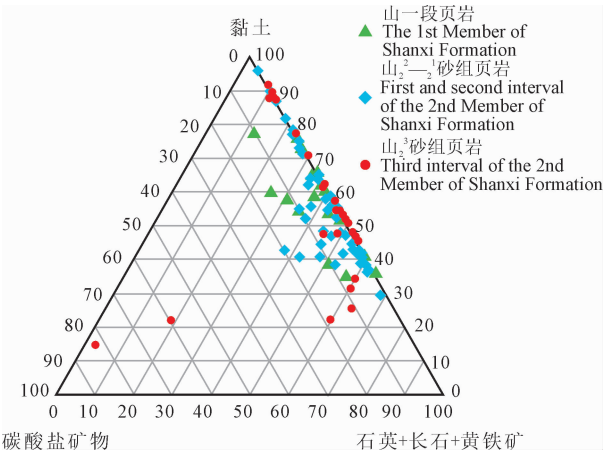


图 7 鄂尔多斯盆地山西组页岩矿物组成三角图
Fig. 7 Triangle diagram of mineral composition of Shanxi shale in eastern Ordos basin

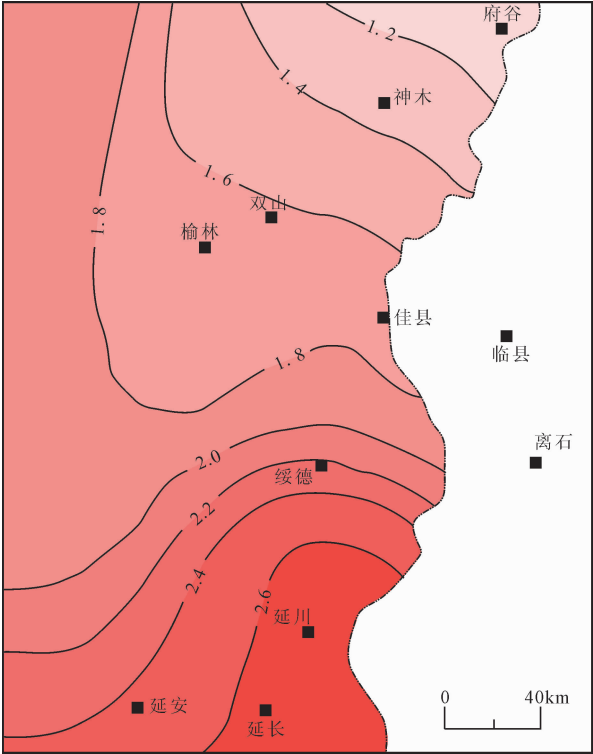


图 8 鄂尔多斯盆地东部山西组页岩 $R_c(\%)$ 等值线图
Fig. 8 $R_c(\%)$ isopach of Shanxi shales in eastern Ordos basin

这些原始有机质易于形成 I、II 型为主的干酪根,在有机质成熟过程中会先生成大量液态烃,而这些液态烃在温度继续升高情况下会裂解形成甲烷,同时形成发育孔隙的沥青残余物。并且,海相页岩多形成于闭塞海盆底部,缺氧环境使硫还原进入沉积物形成硫化物,也有利于有机分子成环和芳构化,热演化过程中容易形成沥青,从而保存气泡成因的纳米有机质孔隙(Schenk et al., 1997)。

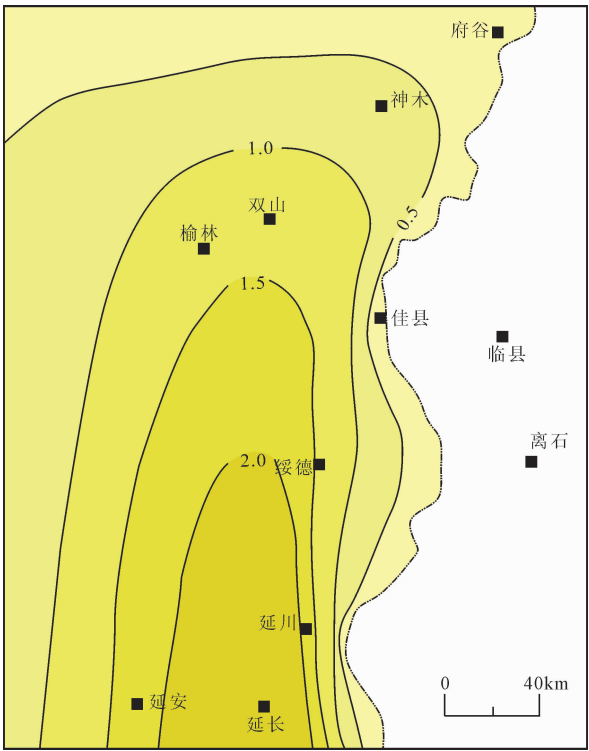


图 9 鄂尔多斯盆地东部山西组页岩含气量(m^3/t)等值线图
Fig. 9 Gas content (m^3/t) isopach of Shanxi shales in eastern Ordos basin

研究区山西组页岩主要形成在三角洲平原的沼泽环境和三角洲前缘分流间湾环境(Guo Wei et al., 2015)。页岩中形成有机质的原始物质主要为高等植物或湖泊水生生物,富含纤维素和木质素、蛋白质、类脂化合物,含有相对高分子量正构烷烃($\text{C}_{23} \sim \text{C}_{33}$)、各种萜类和藿烷甾族化合物(C_{27} 、 C_{28} 和 C_{29})、类异戊二烯烃(Schenk et al., 1997; Zhao Dife et al., 2018)。这些原始有机质形成的干酪根类型主要为 III 型(Wang Shejiao et al., 2011; Zou Caineng et al., 2011; Zhang Lixia et al., 2012; Fu Jinhua et al., 2013; Yan Deyu et al., 2013; Yang Hua et al., 2013)。这使得山西组页岩有机质在成熟度逐渐升高过程中大多数直接生成甲烷,液态烃产量很少,没有发生大量的液态烃裂解生成甲烷,因此缺乏大量的液态烃裂解残余产物(沥青)及其中发育的气泡孔隙。这与观察到的山西组页岩孔隙中有机孔占比小于 30% 的现象相吻合。

4.2 山西组不同岩性空间配置关系及其对页岩气成藏的制约

目前已商业开发的海相页岩,包括美国的海相页岩和中国四川盆地志留系龙马溪组海相页岩,其连续厚度普遍较大,一般超过 30m(Nie Haikuan et

al.,2009; Wang Shiqian et al.,2009; Zou Caineng et al.,2011; Li Jianzhong et al.,2012)。海相页岩成熟过程中干酪根会先生成大量液态烃,但是较大的厚度不利于液态烃的排出,从而使较大数量的液态烃保留其中。继续升温过程中,液态烃裂解生成气态烃和富孔隙的残余沥青,并且生成的气态烃也由于海相页岩较大的厚度而不易排出,致使海相页岩往往含气量较高,包括吸附在残余沥青内的气态烃(图 10a)。

研究区位于鄂尔多斯盆地东部,山西组沉积时期位于晚古生代大华北克拉通盆地中部,物源主要来自盆地北部,沉积环境以三角洲-湖泊相为主(例如, Yu Xinghe et al.,2017; Chen Hongde et al.,2011; Guo Wei et al.,2015; Guo Yinghai and Liu Huanjie,1999),岩性主要由页岩和砂岩组成,下部的山二段发育 3~5 层煤。砂岩和页岩单层厚度较薄,且横向分布不稳定,厚度变化较大;煤层相对稳定,但是厚度变化也往往较大,多分布在页岩顶部。总体上,砂岩与页岩-煤组合呈互层关系,形成了具有“三明治”式特征的岩性空间结构(图 10b)。

山西组这种特殊的岩性结构以及以Ⅲ型干酪根为主的有机质类型,共同导致了山西组页岩或煤生成的天然气能够近距离、有效地排出到相邻砂岩中(即高效地排烃效率),形成常规砂岩气藏或致密砂岩气藏,从而导致页岩或煤中残余的天然气含量相对较少(图 10b),这可能是山西组页岩含气量较低的重要机制,也是与海相页岩气藏成藏特征的显著区别。

5 山西组页岩气藏勘探策略

勘探策略的制定与油气藏储量丰度关系密切。若某类岩性的油气藏储量丰度较高,可采取单独开采的策略。但是,当某类岩性的油气藏储量丰度低到实施单独开采无法获得经济效益时,则必须考虑其它勘探策略,以便于经济动用这类油气藏。

研究区山西组页岩含气量较低,单层厚度薄,横向分布不稳定,以及特殊的成藏条件都决定了山西组页岩具有较低储量丰度。因此,单独开采山西组页岩很难获得较好的经济效益。并且,山西组页岩脆性矿物含量较低,意味着储层改造效果也将受到一定程度影响,这进一步加剧了单独开发山西组页岩气藏的风险。对数口井页岩的试气效果证实了山西组页岩普遍较低的储量丰度。另一方面,研究区山西组的致密砂岩中往往存在具有较好物性的储

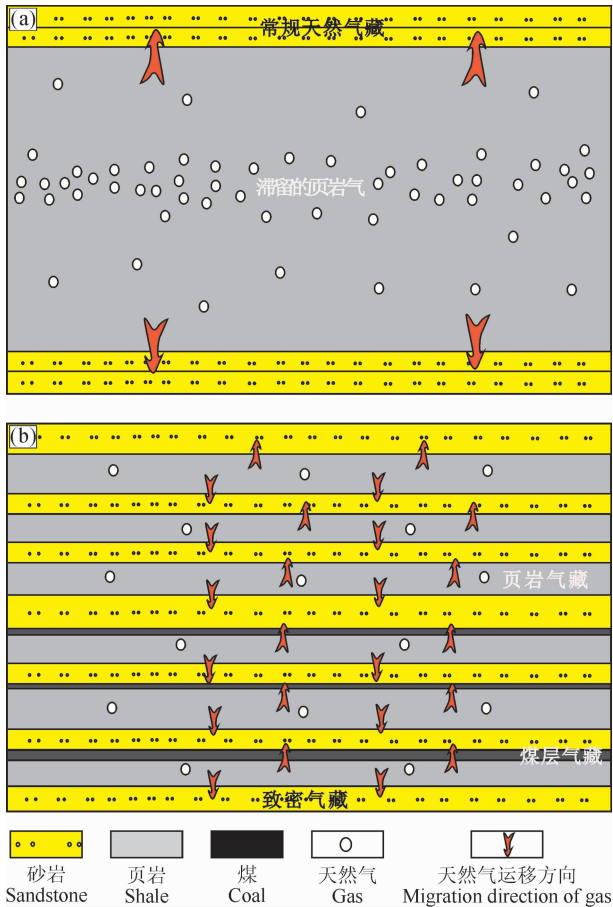


图 10 典型海相页岩气(a)与山西组页岩气(b)成藏模式对比
Fig. 10 Comparison of shale gas accumulation modes between typical marine shale (a) and Shanxi shale (b)

层,含气量高,这部分储层已获得商业开发,如子洲气田及其以北地区的部分山西组储层。但是其余的物性相对较差部分砂岩因为相对较为致密和含气量较低而未动用。此外,页岩顶部的煤层普遍含气量较高,一般可达到 6m³/t,但是单层厚度往往小于 4m,难以单独开发。

由此,可以考虑对研究区山西组页岩气藏的勘探开发采取立体勘探开发的策略,即在页岩厚度大、含气量高的有利区内,将致密砂岩、页岩和煤层划分为一个开发层系,针对不同岩性的储层采取不同的增产工艺,最后合层开采,以获得最大的单井产量,实现山西组天然气的最有效动用。美国皮渗斯盆地的白河隆起气田地质特征与鄂尔多斯盆地类似,存在煤层气、页岩气和致密砂岩气纵向互层,采用立体勘探开发十分成功(Olson,2003)。因此,在鄂尔多斯盆地东部地区,山西组气藏在勘探开发上应加强煤层、页岩与致密砂岩气藏的综合评价,实施以“纵向含气带”为目标的煤系气立体勘探开发思路,同时

借鉴页岩气水平井分段体积改造工艺技术,最大程度地改造储层,才能获得较高的单井产气量,实现山西组页岩气藏的经济开发。

6 结论与认识

(1)鄂尔多斯盆地东部山西组沉积环境主要为近海沼泽-三角洲沉积环境,山西组发育黑色页岩、碳质页岩夹泥质粉砂岩及煤层,页岩具有厚度大、分布广泛、总有机碳含量高、演化程度高特征,具备页岩气形成的基本地质条件。

(2)鄂尔多斯盆地东部山西组页岩干酪根类型主要为Ⅲ型,液态烃产量很少,难以形成大量固化的有机质纳米气泡孔,储集空间类型中有机质纳米孔占比较少,有机质中含气量低于海相页岩。

(3)鄂尔多斯盆地东部山西组页岩气成藏存在特殊性,有机质纳米孔不发育,砂泥煤互层发育,大量页岩气已运移到附近砂岩中形成致密气,煤层、页岩和致密砂岩呈三明治式成藏,东南部为成藏有利区域,在勘探页岩气的同时考虑煤系“含气带”立体勘探寻找有利的煤系气富气层段。

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Geological characteristics and exploration countermeasures of shale gas in the Shanxi Formation of the Ordos basin

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Abstract

Shale gas resources of the transitional facies account for about one-third of the total amount in China, and have broad prospects for exploration and development. A breakthrough in exploration and development has not been achieved despite the large thickness and wide distribution of the transitional facies in the Ordos basin. This study aims to determine differences in the geological characteristics between the transitional and marine shale gas reservoirs and their reservoir-forming principles by analyzing the total carbon content, vitrinite reflectance, mineral composition, micropore texture and gas content in the eastern Ordos basin. The comparative study of shale gas generation, storage and accumulation processes between transitional and marine shale gas reservoirs shows that the shale reservoir formation process of Shanxi Formation has its particularity. Its kerogen belongs to type III originating from higher plants, with little liquid hydrocarbon production, which is bubble pore formation process with a small amount of organic nanopore. Interbeds of sand-mud-coal in the formation mean that the source rock has a high efficiency of hydrocarbon expulsion, resulting in low content in the shale. The shale gas accumulation process of the Shanxi Formation is quite different from that of the marine shale gas because of its high expulsion efficiency and influence on shale gas content. After generation within the Shanxi Formation shale, a lot of gas has migrated to the overlying or underlying tight sandstones. The tight sandstones, coals and shale reservoirs resemble a sandwich. It is predicted that the southeast of the study area is a favorable area. When exploring the shale gas in Shanxi Formation, the “gas bearing zone” of coal series should be comprehensively considered, which means favorable shale gas-rich zones should be investigated together with tight sandstones and coals using three-dimensional exploration strategy for commercial production.

Key words: Ordos basin; Shanxi Formation; shale gas; organic matter nanopore